

## Aerosol size distribution retrieved from optical depth measurements in Tamanrasset and Blida

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**Abstract** - *Since we are often faced with limited, insufficient or contaminated observations in remote sensing. We will employ a theoretical technique for the retrieval of the particle size distribution. In this study, the columnar aerosol size distributions have been inferred by numerically inverting particulate optical depth measurements obtained in Blida and Tamanrasset from AERONET database at wavelengths region between 0.340 - 1.640  $\mu\text{m}$ . Our results are in good agreement with those found by other authors and the correlation between the aerosol optical thickness measurements and the reproduced AOTs by the inverted columnar aerosol size distribution is very good with a correlation coefficient more than 0.97.*

**Résumé** - *Car nous sommes souvent confrontés à des observations limitées, insuffisantes ou contaminées dans la télédétection. Nous allons employer une technique théorique pour l'extraction de la distribution granulométrique des particules. Dans cette étude, les distributions en taille des aérosols intégrés sur la colonne atmosphérique ont été obtenues par l'inversion numérique des mesures de l'épaisseur optique des particules établis à Blida et Tamanrasset à partir des données d'AERONET dans la région spectrale entre 0.340 - 1.640  $\mu\text{m}$ . Nos résultats sont en bon accord avec ceux trouvés par d'autres auteurs et la corrélation entre les mesures d'épaisseur optique des aérosols et l'AOTs reproduites par la distribution en taille inversé est très bonne avec un coefficient de corrélation plus de 0.97.*

**Keywords:** Aerosol properties - AERONET - Size distribution - Inversion methods.

### 1. INTRODUCTION

Atmospheric aerosols are suspensions of small solid or liquid particles in the atmosphere, which plays an important role in atmospheric and environmental research since they take part in many physical and chemical processes in the atmosphere.

Because of the wide variety of sources, the properties of atmospheric aerosol particles, such as size, shape, chemical composition, and optical thickness, may be heterogeneous, and their temporal and spatial variation can be very large [1-5].

Knowledge of both aerosol optical properties (extinction, scattering cross section, phase function, and single-scattering albedo) and microphysical aerosol properties (particle size distribution function and the complex refractive index) is essential for the determination of the effect of atmospheric aerosols on the climate and for the control of the air quality [6].

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Meanwhile, aerosol particle size distribution and aerosol optical thickness are two important properties for characterizing the aerosol particles and the correction of the atmosphere [7].

While aerosol size distributions can be measured in-situ, in different size ranges, using a variety of instruments [8], they can also be derived from measured aerosol optical properties, such as the spectral aerosol optical depth [9], or backscattering [10], or a combination of angular scattering intensities and spectral optical depth [11, 12].

Various techniques have been developed for recovering the particle size distribution function  $n(r)$ , for example direct regularization methods [7, 13-17] and iterative methods [18-23].

This study deals only with the inversion of spectral aerosol optical thickness measurements, and its aim is to determine aerosol size distributions on Blida (36N, 2E) and Tamanrasset (22.78S, 5.53E) northern and southern Algeria respectively during 2006. To accomplish this, we use the inversion algorithm of King *et al.* [9].

## 2 THE INVERSION PROBLEM

Since the relationship between the size of atmospheric aerosol particles and the wavelength dependence of the extinction coefficient was first suggested by Ångström in 1929, the size distribution began to be retrieved by extinction measurements.

First, Ångström inferred that the parameters of a Junge size distribution could be obtained from the aerosol optical thickness (AOT) at multiple wavelengths, and he obtained the useful Ångström empirical formula of Junge size distribution  $\tau_{ae} = \beta \times \lambda^{-\alpha}$ , where  $\tau_{ae}$  is the measured AOT,  $\beta$  is the turbidity coefficient, and  $\alpha$  is the Ångström exponent reflecting the aerosol size distribution [24].

Assuming that aerosols are homogenous spheres, the relationship between the aerosol size distribution and AOT  $\tau_{ae}$  can be written as Fredholm integral equations of the first kind,

$$\tau_{ae}(\lambda) = \int_0^{\infty} \pi r^2 \times Q_{ext}(r, \lambda, m) \times n_c(r) \times dr \quad (1)$$

where  $r$  is the particle radius,  $\lambda$  the wavelength of incident illumination,  $m$  the complex refractive index of the aerosol particles,  $Q_{ext}(r, \lambda, m)$ , the extinction efficiency factor from Mie theory, and  $n_c(r)$  the columnar aerosol size distribution, that is, the number of particles per unit area per unit radius interval in a vertical column through the atmosphere.

Various methods that have been proposed for solving this equation include those in which an assumption about the analytical form of the size distribution to be retrieved is made [25, 26], as well as methods that make no assumption about the size distribution shape [9, 27].

## 3 METHOD OF SOLUTION

In this article, the constrained linear inversion is used to solving the equation (1); this method was introduced by Twomey [1] and developed by King *et al.* [9], to

determine  $n_c(r)$  from spectral measurements of  $\tau_{ae}(\lambda)$ , let  $n_c(r) = h(r) \times f(r)$ , where  $h(r)$  is rapidly varying function of  $r$ , while  $f(r)$  is more slowly varying. So, the equation (1) becomes:

$$\begin{aligned} \tau_{ae}(\lambda) &= \int_{r_a}^{r_b} \pi r^2 \times Q_{ext}(r, \lambda, m) \times h(r) \times f(r) \times dr \\ &= \sum_{j=1}^q \int_{r_j}^{r_{j+1}} \pi r^2 \times Q_{ext}(r, \lambda, m) \times h(r) \times f(r) \times dr \end{aligned} \tag{2}$$

Here  $r_1 = r_a$  and  $r_b = r_{q+1}$  and  $r_1, r_2, \dots, r_{q+1}$  are the  $q+1$  boundaries of  $q$  coarse intervals of integration. If  $f(r)$  is constant within each coarse interval. Thus the Fredholm equation is replaced by simultaneous equations

$$\bar{g} = \hat{A} \bar{f} + \bar{\varepsilon} \tag{3}$$

where,

$$\begin{aligned} g_i &= \tau_{ae}(\lambda_i) & i &= 1, 2, \dots, p \\ A_{ij} &= \int_{r_j}^{r_{j+1}} \pi r^2 \times Q_{ext}(r, \lambda_i, m) \times h(r) \times dr & j &= 1, 2, \dots, p \\ f_j &= f(\bar{r}_j) \end{aligned} \tag{4}$$

The elements  $\varepsilon_i$  of the unknown vector  $\bar{\varepsilon}$  represent the deviations between measurement ( $g_i$ ) and theoretical estimate ( $\sum_j A_{ij} \times f_j$ ). These deviations arise from measurement and quadrature (integration) errors and due to uncertainties of the kernel function  $\pi r^2 \times Q_{ext}(r, \lambda, m)$ .

$\bar{r}_j$  are the midpoints of coarse intervals. Note, the solution of equation (3) is obtained in logarithmic scale with respect to particle radius  $r$ .

If  $h(r)$  takes the form of a Junge size distribution [28],

$$h(r) = r^{-(v^*+1)}$$

With this substitution, King *et al.* [9] apply a quadrature method and a minimization procedure leading to the solution vector

$$f = \left( A^T \times C^{-1} \times A + \gamma H \right)^{-1} \times A^T \times C^{-1} \times g \tag{5}$$

where  $C$  is the measurement covariance matrix;  $\gamma$  is a nonnegative Lagrange multiplier (choice of the Lagrange multiplier is discussed by King [29]);  $H$  is a smoothing matrix;  $A$  is the matrix representation of the kernel function.

The smoothing matrix  $H$  is defined by Twomey [30] as:



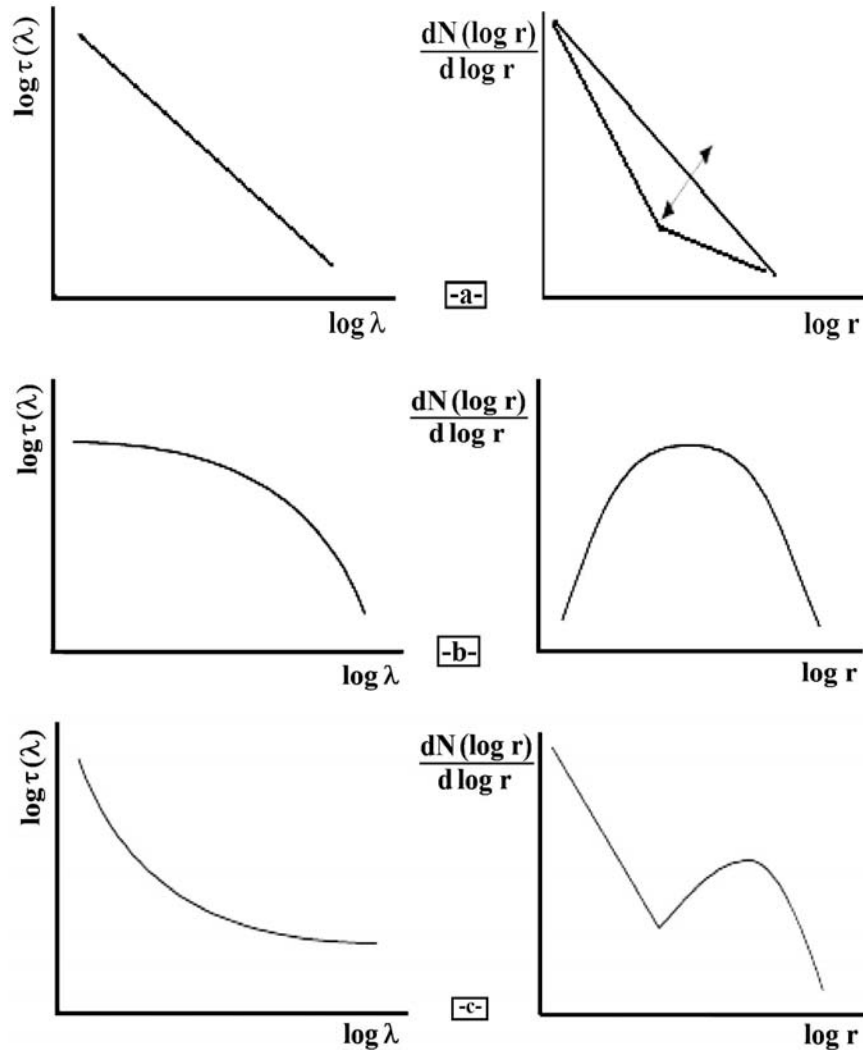


Fig. 1: Measured AOTs and corresponding columnar size distributions  
 -a- Type I; -b- Type II; -c- Type III [9]

## 5. RESULTS

The method for determining the columnar aerosol size distribution  $n_c(r)$  described in the section 3 has been carried out at Tamanrasset and Blida since 2006 from aerosol optical depth measurements by AERONET (AERosol RObotic NETwork) at five and eight different wavelengths ranging between 0.340 and 1.640  $\mu\text{m}$ .

In our study, all inversions were performed assuming the complex refractive index of the aerosol particles was wavelength and size independent and given by  $m = 1.45 - 0.00i$  as in King *et al.* [9].

In lieu of  $n_c(r)$  or, equivalently,  $dN/dr$ , the size distribution results are presented in terms of  $dN/d\log r$ , representing the number of particles per unit area per unit log radius interval in a vertical column through the atmosphere.

### 5.1 In Tamanrasset

From the inspection of results for 192 different days at Tamanrasset, we derive three examples:

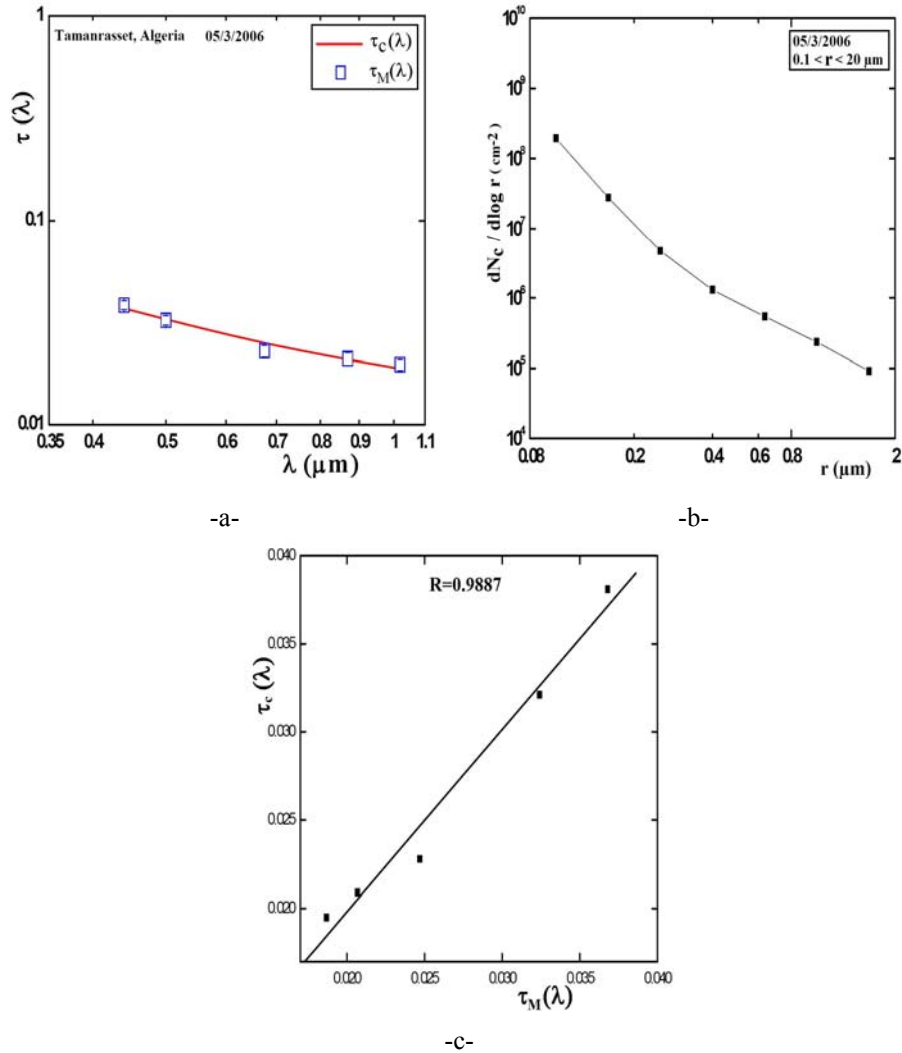


Fig. 2: Measured and retrieved AOT and corresponding columnar size distribution for 05 Mars 2006, (-a-, -b- and -c-)

Fig. 2 illustrates the spectral optical depth measurements for 05 Mars 2006, corresponding size distribution and the linear fit between optical depth measurements

and calculations, this day represents an example of the first type (Type I) for which the measured AOTs nearly follow Ångström's formula given by equation,

$$\tau_M(\lambda) = \beta \times \lambda^{-\alpha}$$

The observed Mie optical depths and corresponding standard deviations are shown in the left portion of the figure while the size distributions obtained by inverting these data is shown in the right portion.

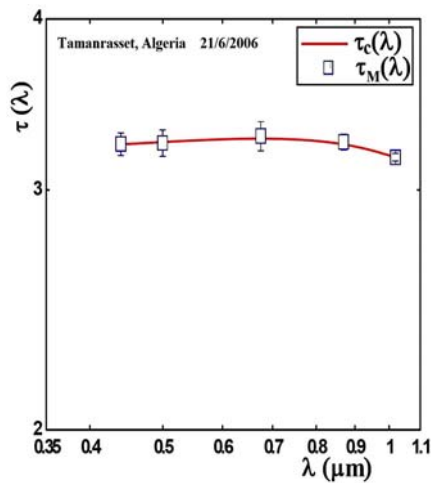
$\tau_c(\lambda)$  in the left portion indicates how the inverted size distributions are able to reproduce the AOTs measurements (i.e., the direct problem  $g = A \times f$ ). The aerosol size distributions illustrated in the same figure can be best described as constructing Junge or two slopes type of distributions.

On the same figure, we consider the linear fit between  $\tau_M(\lambda)$  and  $\tau_c(\lambda)$ , we see that the correlation is good with a correlation coefficient  $R = 0.9887$ .

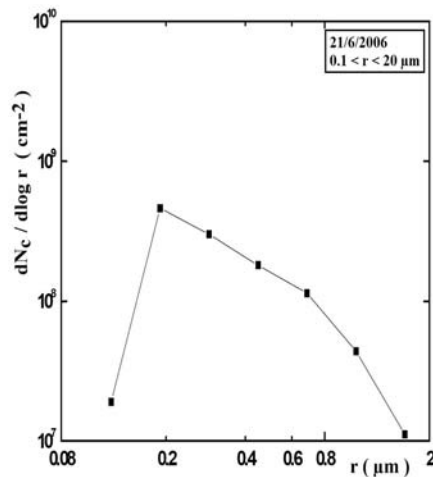
For the data of 21 June, when measured AOTs exhibit small negative curvature (Fig.3) the solutions tend to be monodisperse (Type II). This is not unexpected because the tendency for negative curvature suggests an absence of both small and large particles.

The optical depth measurements which produce this second class of distributions are typically very large. An important modification is the case when experimental data constitute two overlapping monodisperse distributions (we can see this case on Blida results presented below).

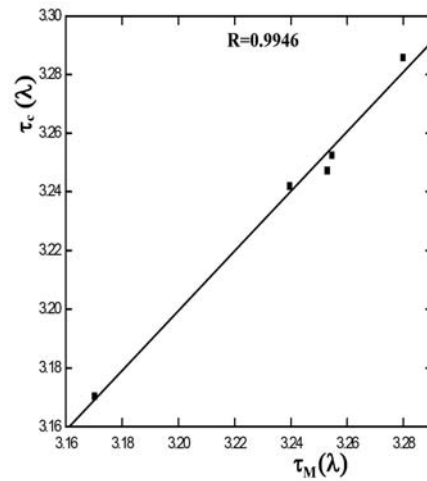
For the correlation between the observed Mie optical depths  $\tau_M(\lambda)$  and the reproduced AOTs  $\tau_c(\lambda)$  is always good and  $R = 0.9946$ .



-a-



-b-

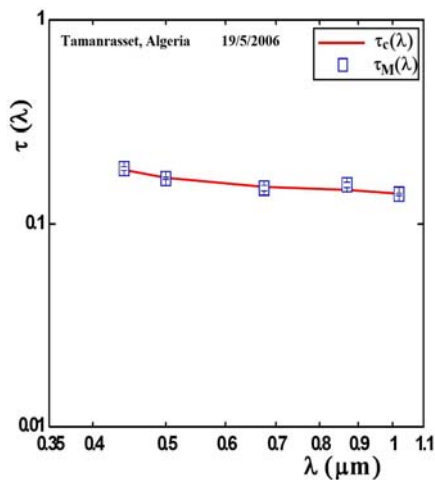


-c-

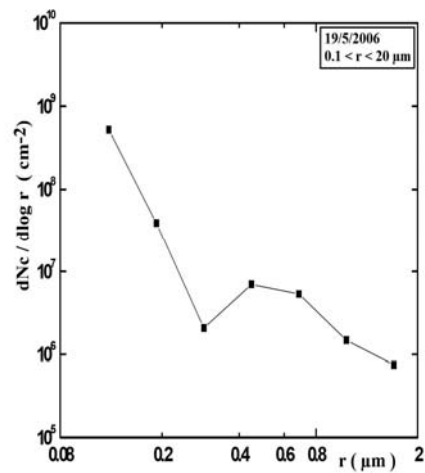
Fig. 3: Measured and retrieved AOT and corresponding columnar size distribution for 21 June 2006, (-a-, -b- and -c-)

The most interesting distribution type is one for which the AOTs intermediate between these of type I and type II. In this case  $\tau_M(\lambda)$  tends generally to have positive curvature.

An example of this type (Type III) is illustrated in Fig. 4 for the data of 19 May. The solution is usually a bi-modal distribution and the correlation coefficient between the observed and the reproduced AOTs is  $R = 0.9700$ .

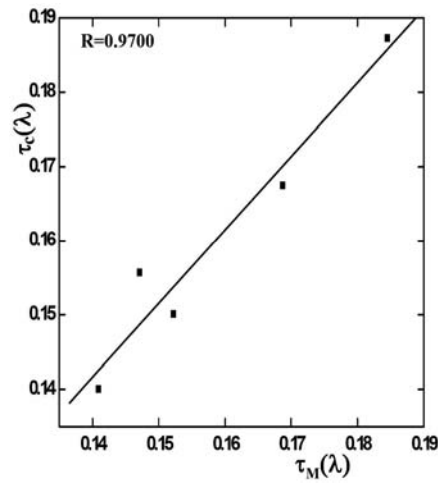


-a-



-b-





-c-

Fig. 4: Measured and retrieved AOT and corresponding columnar size distribution for 19 May 2006, (-a-, -b- and -c-)

### 5.2 In Blida

The inspection of plots of  $\log \tau(\lambda)$  vs.  $\log \tau_M(\lambda)$  (Fig. 5) reveals the presence of at least two kinds of aerosols with possible bi-modal and Junge or tow slopes distributions because plots on May to November have positive curvature and all other curves nearly follow Ångström’s formula. The retrieved columnar number distributions (Fig. 5) are quite interesting.

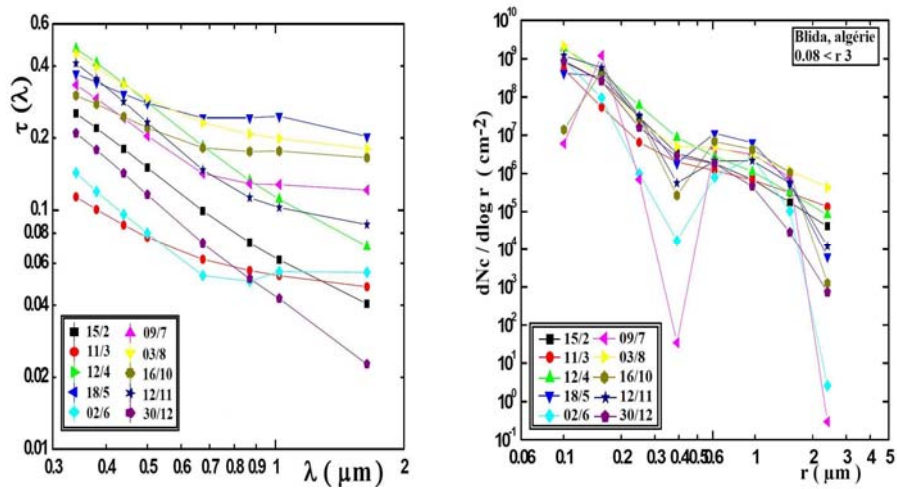


Fig. 5: Measured and retrieved AOT and corresponding columnar size distribution during 2006 in Blida

For 15 February, 11 March, 12 April and 30 December are constituted by aerosols with quite similar properties and distributions are Junge-type. The 18 May, 02 June, 09 July, 16 October and 12 November aerosols ensemble has all bi-modal distributions with a large amount of small particles, deep minimum at  $0.3870 \mu\text{m}$  and maximum at  $0.9718 \mu\text{m}$ .

Distributions of 03 August have similar but smoother shape – minimum at  $0.3870 \mu\text{m}$  is much shallower and about six orders more particles are present at largest radii compared with 09 July.

The data case of  $\tau(\lambda) M$  for 18 May has small negative curvature at the four longest wavelengths, and the corresponding size distribution could equally well be categorized as a type II (monodisperse) distribution composed by two monodisperse which could be fine ash and water droplets, thus making classification according to three distinct types somewhat arbitrary [9].

## 6. CONCLUSION

In this paper, we investigate linear optimization method for the solution of the atmospheric aerosol particle size distribution function retrieval problem from spectral measurements of the particulate (Mie) optical depth at wavelengths region between  $0.340$  and  $1.640 \mu\text{m}$ .

We first described this method of solution developed by King in 1978 and classified experimental data and expected solution type. Then, we applied the proposed method on Tamanrasset site southern Algeria for 192 days during 2006.

When, the columnar aerosol size distributions have been determined from data delivered by AERONET network at five wavelengths  $0.440$ ,  $0.500$ ,  $0.675$ ,  $0.870$  and  $1.020 \mu\text{m}$ . The optical depth measurements and corresponding aerosol size distributions are illustrated in Fig. 2-4 for a few of these days. Our results are in agreements with those found by King in Tucson, Arizona.

His results generally indicate (at least for  $r \geq 0.1 \mu\text{m}$ ) that the aerosol size distribution on a particular day can be represented either as a Junge distribution (type I), a relatively monodisperse distribution such as a log-normal or gamma distribution (type II), or as a two-component system consisting of a combination of both of these types (type III).

Finally, we want to apply the method on Blida site northern Algeria from spectral data at eight wavelengths  $0.340$ ,  $0.380$ ,  $0.440$ ,  $0.500$ ,  $0.675$ ,  $0.870$ ,  $1.020$ ,  $1.640 \mu\text{m}$ , when the retrieved columnar number distributions are quite interesting, what is represented on Fig. 5.

The presence of at least two kinds of aerosols is very clear with bi-modal (Type III) and Junge-type (Type I) distributions. The type I distributions occur mainly in the winter and early spring months and the type III occur throughout the late spring to late summer months.

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